TURBULENT SIMULATION OF FLOW IN OPEN CHANNEL WITH ONE FLOODPLAIN AT Re = 15000

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ABSTRACT

With the avaliability of modem technology, it is now possible to integrate numerically the Navier-Stokes equations and simulate three dimensional flow in open channels. A numerical technique for simulating turbulent flows in an open channel with one flood plain at Reynolds number of 15000 has been applied. The results were good and showed eddy existence and separation.

Keywords: Compound channel flow, turbulent channel flow, floodplaim, flow at low Reynolds number computational modelling.

1. INTRODUCTION

The turbulent motion of a liquid at a surface is much less understood than the corresponding motions at a solid boundary, but the practical applications are equally important. For example, the dispersion of pollutants in rivers and coastal waters is governed by surface phenomena. It is natural to turn to direct numerical simulation of the turbulence, and this approach is able to produce data that is more complete and often more accurate than experiemnts for near wall boundary layers but is limited to low Reynolds numbers. The Large Eddy Simulation approach is closely related to direct numerical simulation but rotation only the large eddies or grid scales with the smaller eddies or subgrid scales being represented by a subgrid model. The method thus has the disadvantage of being dependent on a model but experience has shown that the dependence is much less severe than for other methods such as k-e or algebraic stress models. The large eddy simulation approach is able to extend the Reynolds number of a simulation without increasing the computational resources required. A review of numerical simulation is given by Rogallo & Moin (1984), and some recent simulations of channel flow are given by Kim, Moin & Moser (1985) and Ral & Moin (1991).

2. COMPUTATIONAL DOMAIN

We consider the turbulent flow of an incompressible fluid of kinematic viscosity v maintained by gravity in an infinite open channel of depth d. The cartesian coordinates (x,y,z) are aligned with the channel and dimensions as shown in Figure (1). The channel slopes at a small angle a relative to the horizntal so that the flow is maintained with mean bed stress $\tau_b = \frac{1}{2} U_\tau^2$ at the floodplain, whereas $\tau_b \sim 1.3 U_\tau^2$ at the channel bottom, where U_r denotes the characteristic shear velocity $(\alpha Rg)^{\frac{1}{2}}$, and g denotes the acceleration due to gravity.

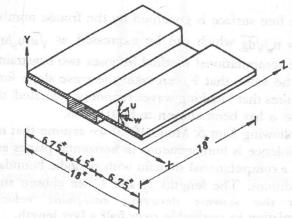


Figure 1. Channel dimensions.

The two characteristic lengths for this flow are the depth and the viscous length v/U_r which governs the structure of the flow near the bed.

The Reynolds number $Re = 4R \text{ ub/}\tau$ based on average velocity over the area = 1ft./sec. = 30.48 cm/sec., and the channel depth is given by (U_b/U_c) Re where U_b/U_r is determined from the simulation $U_r = 0.011$ m/s., $h^+ = hu_r/v$, $\Delta x^+ = \Delta x U_r/v$. The mesh distribution has been chosen such that: 15 elements in uniform distribution in the floodplain y(1)=...=y(15), and 35 elements of non-uniform distribution such that y(n+1)=y(n)* const. starting from y(16). In the horizontal direction, the domain is devided uniformly x(n+1)=x(n)...[n=1.64]. In the Zdirection the domain has been divided into 96 elements in each side of the floodplain, and 63 elements in the channel. Figure (2) outlines the channel cross-section with the mech distribution such as 64 50 255.

$$\Delta x = \frac{18}{64} = 0.2813'' = 0.0117' = 3.572 \times 10^{-3} \text{m}$$

$$\Delta_{x}^{*} = \frac{3.572 \times 10^{-3} \times 0.011}{1 \times 10^{-6}}$$

$$\Delta y_f = \frac{0.25}{15} = 1.67 \times 10^{-2} \text{in.} = 39.3 \text{ x} \frac{1.67 \times 10^{-2}}{0.2813} = 2.3^{+2}$$

$$\Delta y_{m/c} = \frac{1.75}{35} = 0.05 \text{in} = 39.3 \times \frac{0.05}{0.2813} = 7.0^{+2}$$

$$\Delta_z = \frac{18}{255} = 7.06 \times 10^{-2} \text{in} = 39.3 \times \frac{7.06 \times 10^{-2}}{0.2813} = 9.9^{+}$$

The free surface is governed by the froude number $F_r = u_b/\sqrt{dg}$ which can be expressed as $\sqrt{\alpha}(u_b/u_r)$. The computational method imposes two constraints on the value that F_r can take: the wave slope limit requires that breaking waves cannot be tolerated, the slope α has been chosen to be 0.00085.

Following Kim & Moin (1985) we assume that the turbulence is homogeneous in horizontal planes and use a computational domain with periodic boundary conditions. The lengths I_z and I_v are chosen such that the slowest decaying two-point velocity correlation in negligible over half a box length.

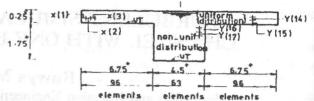


Figure 2. Channel croos-section and mech distribution.

3. NUMERICAL TECHNIQUE

The velocity field $U_1 = (u,v,w)$ satisfies a discretized form of the Navier-Stokes equations and the surface elevation h satisfies $\partial h/\partial t = G$ where G is defined below. We use second order finite differences on a staggered mesh with an unstructured volume of fluid type treatment at the free surface. The convection terms are approximated by central difference type operators which conserve total momentum and kinetic energy in the presence of arbitary surface deformations.

$$G_i^n = \left[-\int_o^h (\frac{\delta u}{\delta x} + \frac{\delta v}{\delta y}) dz \right]^n$$

$$H_i^n = \left[-\frac{\delta u}{\delta x_j} + vD^2 u_i + g_i \right]^n$$

where n denotes the discrete time, level, δ the finite difference operator, D^2 the discrete Laplacian. The time advancement scheme used originally was only first order accurate and although it conserved the potential energy of surface waves it proved to be unsuitable for simulating turbulence. We require that the time truncation error associated with the convection terms be very small compared with the viscous dissipation term, and this could be satisfied only for unacceptably small time steps.

$$\frac{\hat{u}_{i} - u_{i}^{n}}{\Delta t} = \frac{3}{2} H_{i}^{n} - \frac{1}{2} H_{i}^{n-1} + \frac{1}{2} \frac{\delta p^{n-1}}{\delta x_{i}}$$

$$\frac{u_{i}^{n+1} - \hat{u}_{i}}{\delta t} = -\frac{3\delta p^{n}}{2\delta x_{i}}$$

$$D^{2} p^{n} = \frac{2\delta \hat{u}_{j}}{3\Delta t \delta x_{i}}$$
(1)

 $\frac{h^{n+1} - h^n}{\Lambda_t} = \frac{1}{2}G^{n+1} + \frac{1}{2}G^n$

where $\hat{\mathbf{u}}_i$ denotes an intermediate variable. The continuity equation $\delta \mathbf{u} \mathbf{j}^{n+1}/\delta \mathbf{x}_j = 0$ is enforced at time level n+1. First, $\hat{\mathbf{u}}_i$ is computed using (1) and them \mathbf{p}^n is determined such that \mathbf{u}_1^{n+1} satisfies continuity by solving (3). The boundary conditions are assumed built into the difference operators and need not be explicitly treated. The thickness of the layer is an integral measure of the eddy size.

4. MODEL RESULTS

After running the model with respect to time, it has been found that eddies and separation at the floodplain and channel intersection, also velocity vectors had apparent values in all directions.

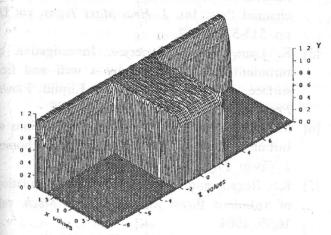


Figure 3. Velocity distribution after 4000 time steps.

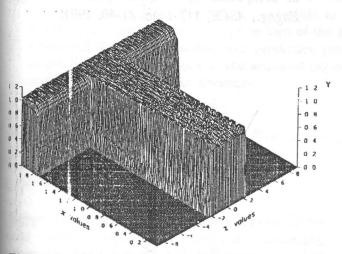


Figure 4. Velocity distribution after 8000 time steps.

Figure (3) shows the velocity distribution after 4000 time step, while Figure (4) shows the velocities after 8000 time steps showing bending of the profile in the channel bottom. Figure (5) outlines definite velocities and some separation along the intersection after 20,000 time step, while Figure (6) describes this situation in a cross-section. Finally, after 40,000 time steps, Figure (7) shows the final profile with eddy separation and velocities which is also decsribed in cross-section profile in Figure (8).

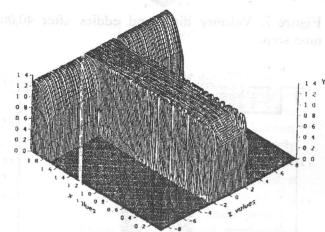


Figure 5. Velocity distribution after 20,000 time steps.

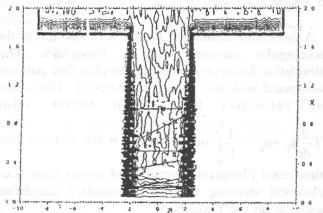


Figure 6. Cross-section showing the distribution after 20,000 time step.

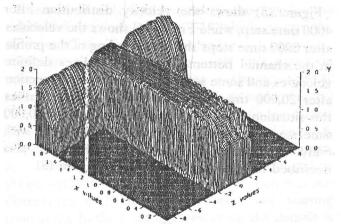


Figure 7. Velcoity distr. and eddies after 40,000 time step.

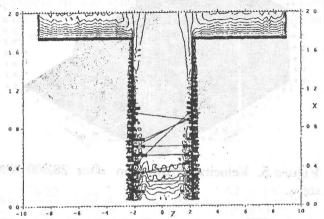


Figure 8. Final vel. distr. and separtation in the cross-section.

5. CONCLUSION

A successful simulation has been carried out of the rectangular channel with a floodplain. This simulation has reproduced the complex flow patterns associated with this particular geometry. The values of velocities have been found using

 $T\frac{d}{dt}u_r + u_r = \frac{1}{L}\int_0^L u dx$ where T is the characteristic

time used fllodplain, overall good results have been obtained showing extended secondary circulation existing on the floodplain.

6. REFERENCES

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